

INTERNATIONAL ENERGY AGENCY

Implementing Agreement for Co-operation in the Research, Development and Deployment of Wind Turbine Systems Task 11

# 51<sup>st</sup> IEA Topical Expert Meeting

# State of the art of Remote Wind Speed Sensing Techniques using Sodar, Lidar and Satellites

Risø, Roskilde, Denmark, January 2007 Organised by: Risø







Scientific Co-ordination: Sven-Erik Thor Vattenfall AB, 162 87 Stockholm, Sweden

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Topical Expert Meeting #51

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#### **INTRODUCTORY NOTE**

#### IEA TOPICAL EXPERT MEETING 51

ON

#### STATE OF THE ART OF REMOTE WIND SPEED SENSING TECHNIQUES USING SODAR, LIDAR AND SATELLITES

Ioannis Antoniou, Torben Mikkelsen, Hans E. Jørgensen, Charlotte Bay Hasager, Jakob Mann

Risø National Laboratory

#### Background

Wind power is moving towards the installation of wind farms in complex terrains, off-shore, in forests, and at high levels in the atmosphere. Marketing of large, multi-MW wind turbines is in continued growth. At the same time our basic knowledge on winds in these challenging environments is inadequate.

The method traditionally used for accredited measurements for wind energy purposes is to mount cup anemometers on met masts. As turbines grow in height, mast instrumentation, erection and maintenance, has become expensive; prices increase geometrically with height and built permits can be time consuming. At the same time the discrepancies between the measured wind at the rotor centre and the turbine performance have increased the need for knowing and measuring the wind over the whole turbine rotor.

Successful development of wind power should be based on sound information on winds in each location. To achieve this it is relevant to place emphasis on new observation methods and strategies. Most promising are the new (for wind energy purposes) remote sensing techniques SODAR, LIDAR and satellite. SODAR is based on sound propagation, LIDAR on laser Doppler and satellite on microwave scatterometry and Synthetic Aperture Radar (SAR) methods. Advantages and limitations of the various techniques will be described and discussed.

#### Techniques

Briefly described the SODAR, LIDAR and satellite techniques for wind observation are summarized below:

SODAR (SOund Detection And Ranging) provides a method for wind speed measurements. The instrument is ground based and emits a short pulse of sound at a certain frequency to the atmosphere. The sound propagates upwards while at the same time a part of the sound is reflected back. The Doppler frequency shift of the received signal is proportional to the wind speed aligned to the transmission sound path. By combining three or five of these pulses, usually one along the vertical and two or four inclined to the vertical, the three dimensional velocity field of both the mean values and the turbulent values is calculated.

SODARs are widely used for meteorology applications however their usage in wind energy, e.g. for measuring the wind field or the energy potential at a site or power curve measurements, is relatively new and involves a number of advantages and drawbacks. Among the advantages, the SODAR gives the possibility to measure the wind profile over the whole rotor, it is ground based instrument and therefore it is faster, easier and cheaper to use relative to cup anemometers mounted on met masts. Among the drawbacks, the most serious are the limited experience in the use of the instrument, its decreasing performance with height, its dependence on the prevailing atmospheric conditions and finally the need for a rigorous wellestablished "absolute" calibration method. Among the SODAR users, there is also a debate as to what degree the instrument can be used for the measurement of turbulent quantities other than the one in a vertical direction and still there is an open question to what extend can the instrument be used for measurements in complex terrain as the separate wind components are not being estimated within the same volume.

LIDAR is a remote sensing technique that offers the ability to determine wind speed and direction at substantial heights using a ground-based instrument. In this respect it is similar to SODAR but operates via the transmission and detection of light rather than sound. The basic LIDAR principle relies on measuring the Doppler shift of radiation scattered by natural aerosols carried by the wind. Typically, these are dust, water droplets, pollution, pollen or salt crystals. A new generation of fibre-based LIDARs has emerged the recent years that operates close to the theoretical limit of sensitivity and typically only needs to detect one photon for every 10E+12 transmitted in order to measure wind speed. As the Doppler-shifted frequency is directly proportional to line-of-sight velocity, the wind speeds obtained by LIDAR instrument seem not to need calibration. This however remains still to become documented by more measurements and by a full description of the whole measurement chain. As in the case of SODARs, the LIDAR is also a new instrument and its merits and limitations are neither fully documented nor are they known. In the case of the LIDAR, the measurement of the wind speed takes place on the surface of a cone where the depth changes as a function of the focus distance. The measurement of the turbulence quantities using LIDARs remains also to be documented.

Satellite remote sensing provides wind maps (snap-shot images) of the surface wind at 10 m above sea level. From scatterometer twice-daily wind maps at grid resolution of 25 km are available. The data series from July 1999 to present holds more than 5000 observations at most locations of the globe. Due to the resolution of 25 km observations are not available close to the coastline (usually a void around 40 to 50 km distance offshore). In contrast, SAR wind maps are available (e.g. a few hundred or less), but using statistical treatment of few samples, rough estimates of the wind resource can be obtained. The accuracy, around 1.1 m/s standard error on a series of wind maps compared to offshore mast observations is useful in pre-feasibility and for decision on siting of offshore masts (or LIDAR/SODAR). In addition, if high-quality met-observations are available within a mapped area, the relative differences in winds between different locations can be estimated with higher accuracy, possibly around 0.6 m/s.

#### TENTATIVE AGENDA

The tentative agenda covers the following items:

- 1. Introduction by host
- 2. Introduction by Operating Agent, Recognition of Participants
- 3. Collecting proposals for presentations. The participants are encouraged to inform the Operating Agent on the contents of their presentation in advance and if possible provide a copy. The participants are also encouraged to in advance suggest relevant discussion matters that would have their interest.
- 4. Presentation of Introductory Note.
- 5. Individual presentations
- 6. Discussion
- 7. Summary of meeting

## Objectives

To hold a symposia meeting to discuss and gather information on:

- Overview of existing knowledge and experience on LIDAR and SODAR technical issues, regarding the measurement of mean wind speeds, turbulence quantities, and vertical wind profiles for wind energy applications.
- Calibration of SODARs and LIDARs.
- Accuracy and reliability of the different systems and comparisons with other point measurement techniques, e.g. cup anemometers
- Suggestion for a "good measurement practice" using remote sensing equipment.
- Overview of existing knowledge on offshore wind mapping from satellite.
- Challenges off-shore compared to on-shore work.
- Getting closer to certification and how?
- Future options for wind energy using LIDAR, SODAR and satellite wind observations.

The participants are encouraged to prepare presentations relevant to these objectives.

## **Expected Outcomes**

One of the goals of the meeting will be to gather the existing knowledge on the subject and come up with suggestions / recommendations on how to proceed with thefollowing:

- 1. Define a procedure of how should the instruments be used in order to make their results acceptable by developers and others active in wind energy?
- 2. How should the instruments be used in different terrain types?
- 3. In what assignments should the instruments be used for (e.g. siting, power curves,...)?
- 4. Limiting the measurement, until further, to only certain parameters (e.g. mean values, turbulence,...)?
- 5. Calibration or verification procedures for the results?

Based on the above a document will be compiled containing:

- Presentations by participants
- Compilation of the most recent information on the topic
- Input to define IEA Wind RD&D's future role in this topic

#### **Intended Audience**

The national members will invite potential participants from research institutions, utilities, manufacturers and any other organizations willing to participate in the meeting by means of presenting proposals, studies, achievements, lessons learned, and others. This means then that the symposia will be wide open, taking into account that it is the first time that this subject will be discussed within the framework of the IEA Wind RD&D.

























































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Overview	
<ul> <li>Technologies Used at NREL         <ul> <li>Satellite wind data (SSMI, QuikScat, etc.)</li> <li>Sodar</li> <li>Lidar</li> <li>SAR</li> </ul> </li> </ul>	
<ul> <li>Priorities</li> </ul>	
<ul> <li>Questions</li> </ul>	
<ul> <li>Future work</li> </ul>	
IEA Expert Meeting on Remote Wind Speed Sensing 4	A NREL National Renewable Energy Laboratory



Satellite Ocean Wind Data		
<ul> <li>Sensors         <ul> <li>Passive (radiometers) – SSM/I, TMI</li> <li>Solve Radiative Transfer Equation</li> <li>Active (scatterometers) – QuikScat</li> <li>Analyze backscattered signa</li> </ul> </li> </ul>		
<ul> <li>Returns wind speed and direction, water vapor and liquid</li> </ul>		
<ul> <li>Accuracy: ±2.0 mps WS, ±20° WD</li> </ul>		
<ul> <li>Less accurate in coastal/shallow regions</li> </ul>		
<ul> <li>RSS daily files combined into monthly 0.25° grids</li> </ul>		
<ul> <li>Monthly grids combined into annual or long-term grids</li> </ul>		
IEA Expert Meeting on Remote Wind Speed Sensing 6		

































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	SGULTENERGY sustainable energy solutions
Vector vs. Scalar averages	
It was observed that although cup anemometers and indeed respond in a way that can be characterised by scalar average $v_s$ , remote sensing devices such as SODAR and LIDAR give from vector averages $v_v$ .	wind turbines d wind speeds results derived
An analytical result describing the ratio of the vector to scalar speed ( $v_v / v_s$ ) that would be obtained for a distribution of wind obtained. This result holds for a constant wind speed and direction distribution of standard deviation $\sigma$ and is shown in f approximations were considered reasonable for deriving a sir result (Bessel function) for 10 minute averages. This result compared to empirical results obtained elsewhere for this ratio.	averaged wind directions was uniform wind igure 3. These mple analytical sult should be
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Outline	
Introduction	
Objectives	
<ul> <li>Participants and funding</li> </ul>	
<ul> <li>System setup (SODAR,power supp., mast, eqipment)</li> </ul>	
Data transmission	
Sensitivity/limitations	
<ul> <li>Results and findings</li> </ul>	
Conclusion	
Recommendations	
References	
Fluid Mechanics Section, MEK Technical University of Denmark	








































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# I) Chances and limitations of measuring wind and turbulence profiles by acoustic remote sensing

## II) Offshore wind and turbulence data

#### Stefan Emeis, Inst. for Meteorology and Climate Research, Dept. Atmospheric Environmental Research (IMK-IFU) Forschungszentrum Karlsruhe GmbH

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IMK-IFU, Stefan Emeis

Forschungszentrum Karlsruhe in der Helmholtz-Gemeinschaft

# I) Chances and limitations of measuring wind and turbulence profiles by acoustic remote sensing

Chances (of profile measurements in general)

#### **Results from the EU-project WISE**

Funded by the European Union under Grant NNE5-2001-297 (partners: ECN, Risø, Univ. of Salford, IMK-IFU, DEWI, Windtest-KWK, CRES)





#### wind speed (scale factor of Weibull distribution)



#### wind variance (shape factor of Weibull distribution)



IMK-IFU, Stefan Emeis



rotor diameter [m] rotor diameter [m]

-1.(-0.50.0 0.5 1.0 1.5 2.0 2.5 3.0 3.5 -1.0 -0.5 0.0 0.5 1.0 1.5 2.0 wind energy error [%] wind speed error [%]

600

wind speed [cm/s], u<sup>a</sup> [m<sup>a</sup>/s<sup>a</sup>]

700 800 1000 1100

1200

900

400

300

IMK-IFU, Stefan Emeis



IMK-IFU, Stefan Emeis

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#### **Observed turning of winds**



## I) Chances and limitations of measuring wind and turbulence profiles by acoustic remote sensing

# Limitations (especially for acoustic remote sensing)

#### **Results from the EU-project WISE**

Funded by the European Union under Grant NNE5-2001-297 (partners: ECN, Risø, Univ. of Salford, IMK-IFU, DEWI, Windtest-KWK, CRES)



IMK-IFU, Stefan Emeis



# II) Offshore wind and turbulence data

# This data is presented here as possible evaluation data for satellite offshore wind mappings

Results from FINO1-measurements (running since Sept. 2003) in the German Bight 45 km off the coast

Funded by the German Ministry for the Environment (BMU) under Grant 0329961 (project: OWID, partners: IMK-IFU, DEWI, DEWI-OCC, GE Wind, Multibrid, Repower, Enercon)



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# FINO1 research plattform

- Measuring of wind components from 33 to 100m
- Monitoring of all standard meteorological parameters
- Measuring of structural loads
- Oceanographic measurements
- Biological measurements
- Located 45km north of the island of Borkum
- ⇒ Long running measurements since September 2003





IMK-IFU, Stefan Emeis





IMK-IFU, Stefan Emeis







IMK-IFU, Stefan Emeis















































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# To extend to calibration Real time recording of pulse Calculation of atmospheric signal based on atmospheric scattering theory and realistic wind profiles Feed signal back into SODAR Compare results with expectations
























## Summary LIDAR and SODAR have generic errors Some errors *can not* be removed via calibration and/or system configuration Bi-static SODAR design offers improved performance through: A single well-defined scattering volume Strong signal dominating all clutter Vertical transmission with reduced side-lobes Current work: auto-alignment & optimized footprint






















































































































Atmospheric inf	fluences			<b>RIS</b> @
- altering the ve	ertical wind p	rofile		
			$\mathbf{p}^2$	
Onshone winds	SD[ms]	Bias [m s ]	<u>R</u>	<u> </u>
Offshore winds	1.10	-0.00	0.89	49
Stable	$\frac{\text{SD} [\text{m s}^{-1}]}{1.47}$	Bias [m s <sup>-1</sup> ] -0.86	$\frac{R^2}{0.88}$	<u>N</u> 11
Near-neutral	0.95	-0.13	0.93	22
Unstable	1.06	-0.26	0.85	52
	SD [m s <sup>-1</sup> ]	Bias [m s <sup>-1</sup> ]	$\mathbb{R}^2$	Ν
	0.93	-0.57	0.90	46
No wind farm		0.04	0.87	45



















DTU	QuikSCAT wind maps for 7 years		RISØ		
<b>₩</b>	Longitude and latitude $^{(\circ)}$	Mean wind speed $_{(m/s)}$	Weibull A (m/s)	Weibull k	
Cape Verde	334.25 E, 16.75 N	8.04	8.86	4.57	
Denmark	7.75 E, 55.50 N	7.95	9.06	2.26	
Cape Horn	67.5 W, 56.50 S	11.2	12.88	2.44	

















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## Other uses of the QinetiQ lidar at Risø

J. Mann, F. Bingöl, G. C. Larsen, E. Dellwik and S. Ott <u>Risø National Laboratory/DTU, Denmark</u>

January 21, 2007

## Other uses of the QinetiQ lidar at Risø

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January 21, 2007

• Flow over a forest: mean wind and turbulence

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January 21, 2007

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- The wake behind a wind turbine

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January 21, 2007

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- The wake behind a wind turbine





Wakes and momentum flux

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Wakes and momentum flux

J. Mann et al.

Because the half opening angle of the cone is  $\approx 30^\circ$  the radial velocity is

$$v_r = \left| \frac{1}{2} u \cos \theta + \frac{1}{2} v \sin \theta + \frac{\sqrt{3}}{2} w \right|,$$

where  $\theta$  is the horizontal angle from the downwind direction. The fluctuations in the upwind ( $\theta = \pi$ ) and downwind ( $\theta = 0$ ) directions are

$$\sigma^{2}(v_{r,up}) = \frac{1}{4}\sigma_{u}^{2} + \frac{3}{4}\sigma_{w}^{2} - \frac{\sqrt{3}}{2}\langle u'w'\rangle$$
  
$$\sigma^{2}(v_{r,down}) = \frac{1}{4}\sigma_{u}^{2} + \frac{3}{4}\sigma_{w}^{2} + \frac{\sqrt{3}}{2}\langle u'w'\rangle$$

so subtracting these equations the momentum flux  $\langle u'w' \rangle$  can be obtained.

```
Wakes and momentum flux
```

Momentum flux profile and average Doppler spectra



4

## The Tellus wake experiment



Wakes and momentum flux

Laser Doppler data (1:5)

### Comparison with averaged wind direction:





### Comparison with averaged wind direction:





#### Comparison with averaged wind direction:







# "TV scanning" of a wake



Wakes and momentum flux

IEA 2007 January Meeting, Risø

J. Mann et al.







Acceptance criteria as the primary wind on the Beatrice Alph	for the ZephIR being used monitioring method a platform:
→ Availability	> 95 % (system & data)
→ Data quality	relative to cups linear regressions through origin
	Y = mx + b (i.e.with b==0)
	0.97 < m < 1
	R <sup>2</sup> > 97%
WINDTEST Kaiser-Wilhelm-Koog GmbH	Acceptance Criteria EA T.E.M. RISØ Jan 2007 Slide No. 4



Data Storage Period No.	Start Date	End Date	Height Settings	Cloud Correction	
1 to 6	14.9.2005	30.9.2005	120 / 300	off	
7 to 16	30.9.2005	8.11.2005	120 / 300	off	
17 to 24	8.11.2005	19.12.2005	60, 90, 120, 150 / 300	off	
24 twin	15.12.2005	19.12.2005	60, 90, 120, 150 / 300	off	
25 to 27	19.12.2005	5.1.2006	60, 90, 120, 150 / 300	on	
Overall S	ystem Availa	bility:	99.6	%	
Overall D	ata Availabili	ity (10-MinA	v.): 95.2	%	
WINDTEST Kaiser-Wilhelm-K	- loog GmbH IEA	Availabili	ty Onshore	Oldbaum Services Ltd.	



Sector	125° to 255°		180° to 255°	
	CUP	CUP	CUP	
1 <sup>st</sup> Period	120 m	90 m	60 m	
10-min-avg. values	3034	1	1	
Slope "m"	0.94	1	1	
Regr. Coeff "R <sup>2</sup> "	0.95	1	1	
	CUP	CUP	CUP	SONIC
2 <sup>nd</sup> Period	120 m	90 m	6	0 m
10-min-avg. values	2532	1688	1577	1568
Slope "m"	0.95	0.97	0.99	1,00
Regr. Coeff "R <sup>2</sup> "	0.96	0.97	0.95	0.93
· · ·			·	







	Period No.	Data Storage Period No.	Start Date	End Date	Heigth Settings	Cloud Correction	
	1	1 & 2	2.3.2006	11.4.2006	78 / 300	on	
	2	3 - 6	11.4.2006	26.6.2006	36, 56, 78, 100 / 300	on	
	2a	7 & 8	26.6.2006	1.7.2006	36, 56, 78, 100 / 300	off	
	2b	9	3.7.2006	5.7.2006	36, 56, 78, 100 / 300	on	
	2c	10	5.7.2006	13.7.2006	36, 56, 78, 100 / 300	off	
	Overall System Availability:				100.0 %		
	Overall Data Availability (10-MinAv.):			<b>99.6</b> %			
	WINDTES	ST		lability Of	fshore oldbo	uum Services Ltd.	








Ottshore					
Analysis Sector	255°, 295°-345°	30° to 90° an	d 180° to 240°	0° to 60° and	1 210° to 270°
. et	400 (70)	CUP	04 (00)		
1 <sup>er</sup> Period	103 (78) m	81 (56) m	61 (36) m		
Olara ""	1905	,	1		
	0.97	1	1		
Regr. Coeff "R-"	0.99		1	02	NIC
and Boried	103 (78) m	81 (56) m	61 (36) m	81 (56) m	61 (36) m
10-min-avg values	6005	2589	2749	3228	3245
Slone "m"	0.98	0.97	0.98	1.01	1.01
Regr. Coeff "R <sup>2</sup> "	0.99	0.99	1.00	0.99	1.000















	125° to 255°		180° to 255°		
No Filtering [a]	CUP	CUP	CUP	SONIC	
2 <sup>nd</sup> Period	120 m	90 m	6	60 m	
10-min-avg. values	2532	1688	1577	1568	
Slope "m"	0.95	0.97	0.99	1,00	
Regr. Coeff "R <sup>2</sup> "	0.96	0.97	0.95	0.93	
Precipitation NO [b]	CUP	CUP	CUP	SONIC	
2 <sup>nd</sup> Period	120 m	90 m	6	60 m	
10-min-avg. values	1787	1209	1146	1133	
Slope "m"	0.95	0.97	0.98	0.99	
Regr. Coeff "R <sup>2</sup> "	0.96	0.97	0.95	0.95	
Precipitation YES [c]	CUP	CUP	CUP	SONIC	
2 <sup>nd</sup> Period	120 m	90 m	6	60 m	
10-min-avg. values	745	479	431	435	
Slope "m"	0.96	0.98	1.00	1.01	
Rear. Coeff "R <sup>2</sup> "	0.96	0.97	0.93	0.89	

















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Solutions under development: corrections, filter

















































### Turbulence measured by the ZephIR<sup>TM</sup>:

# The Effects of Conical scanning and Lorentzian Probe Volume Filtering

by

Torben Mikkelsen and Hans E. Jørgensen

Wind Energy Department Risø National Laboratory

### 1. Introduction

The purpose of this note is to investigate the effective turbulence obtained from horizontal "figure of eight" averaged scans obtained with the conically scanning Coaxial ZephIR<sup>TM</sup> wind Lidar system.

We start with the usual neutral Kaimal spectra:

$$n S_{u}(n)/u_{*}^{2} = \frac{102n}{(1+33n)^{5/3}}$$

$$n S_{v}(n)/u_{*}^{2} = \frac{17n}{(1+9.5n)^{5/3}}$$

$$n S_{w}(n)/u_{*}^{2} = \frac{2.1n}{(1+5.3n^{5/3})}$$
(1.1)

where the dimensionless frequency *n* has been defined as n = fz/U, where *f* is frequency in Hz, *z* is measurement height above the ground, and *U* is the (10-min averaged) mean wind speed.

If we let the upper non-dimensional frequency  $n_{max} = f_{max}z/U$  go to infinity the corresponding definite integrals can be evaluated (u, v analytically, and w via Mathematica , cf. the figures in Eqs.(1.2). For comparison is added: the figures in parentheses () are text book "standard" relations. The figures in (()) come from Panofsky H.A. & J.A Dutton. Atmospheric Turbulence: Models and Methods for Engineering Applications; Wiley: New York (1984).

$$\frac{\sigma_u^2}{u_*^2} = \int_0^\infty \frac{102}{(1+33n)^{5/3}} dn \square 2.15^2; \qquad (2.5^2); \qquad ((2.39^2))$$

$$\frac{\sigma_v^2}{u_*^2} = \int_0^\infty \frac{17}{(1+9.5n)^{5/3}} dn \square 1.64^2; \qquad (2.0^2); \qquad ((1.98^2)) \qquad (1.2)$$

$$\frac{\sigma_w^2}{u_*^2} = \int_0^\infty \frac{2.1}{(1+5.3n^{5/3})} dn \square 1.24^2; \qquad (1.3^2); \qquad ((1.25^2))$$

To investigate the effect of filtering, both by the focussed beams probe volume (the Lorentzian optical probe filter), and by the horizontal conical scanning, we first rewrite the Kaimal spectra in dimensional form:

$$\frac{\sigma_u^2}{u_*^2} = \frac{z}{U} \int_0^{n_{max}U/z} \frac{102}{(1+33fz/U)^{5/3}} df$$

$$\frac{\sigma_v^2}{u_*^2} = \frac{z}{U} \int_0^{n_{max}U/z} \frac{17}{(1+9.5fz/U)^{5/3}} df$$

$$\frac{\sigma_w^2}{u_*^2} = \frac{z}{U} \int_0^{n_{max}U/z} \frac{2.1}{1+5.3(fz/U)^{5/3}} df$$
(1.3)

Eqs.(1.3) is of the form

$$\sigma_i^2 = \int_0^{n_{max}U/z} S_i(f) df;$$

where

$$S_{u}(f) = u_{*}^{2} \frac{z}{U} \frac{102}{(1 + 33fz/U)^{5/3}} \left[\frac{m^{2}}{s^{1}}\right]$$
(1.4)  

$$S_{v}(f) = u_{*}^{2} \frac{z}{U} \frac{17}{(1 + 9.5fz/U)^{5/3}} \left[\frac{m^{2}}{s^{1}}\right]$$
  

$$S_{w}(f) = u_{*}^{2} \frac{z}{U} \frac{2.1}{1 + 5.3(fz/U)^{5/3}} \left[\frac{m^{2}}{s^{1}}\right]$$

To evaluate the effect of spatial filtering we transform these frequency spectra into wave number spectra by use of the relations

$$S_{i}(f)df = F_{i}(k_{1})dk_{1}$$

$$\frac{d\omega}{dk} = U = 2\pi \frac{df}{dk} \text{ (Taylors frozen turbulence hypothesis)}$$

$$\omega = 2\pi f = Uk_{1}$$

$$f = \frac{U}{2\pi}k_{1}; \quad f_{max} = \frac{U}{2\pi}k_{1,max} \Leftrightarrow k_{1,max} = \frac{2\pi}{U}f_{max}$$

$$df = \frac{U}{2\pi}dk_{1}$$

So that

$$\sigma_i^2 = \int_0^{n_{max}U/z} S_i(f) \, df = \int_0^{\frac{2\pi}{z} n_{max}} S_i(\frac{U}{2\pi}k_1) \, \frac{U}{2\pi} dk_1 = \int_0^{k_{1,max}} F_i(k_1) dk_1$$
  
by use of the relations: (1.5)

by use of the relations:

$$f = \frac{U}{2\pi}k_1; \quad f_{max} = \frac{U}{2\pi}k_{1,max}; \quad k_{1,max} = \frac{2\pi}{U}f_{max} = \frac{2\pi}{U}n_{max}U/z = \frac{2\pi}{z}n_{max} \quad ; df = \frac{U}{2\pi}dk_1$$

For instance, the Kaimal u- spectrum in wave number space then looks like:

$$F_{u}(k_{1}) = u_{*}^{2} \frac{z}{U} \frac{U}{2\pi} \frac{102}{(1 + 33\left(\frac{U}{2\pi}k_{1}\right)z/U)^{5/3}} = u_{*}^{2} \frac{z}{2\pi} \frac{102}{(1 + 33\frac{k_{1}z}{2\pi})^{5/3}} \left[\frac{m^{3}}{s^{2}}\right]$$
(1.6)

## 2 Models for the ZephIR<sup>'s</sup> spatial filters due to averaging

#### 2.1 Averaging associated with the Lorentzian optical probe volume:

In the note "On the Lorentzian weighting function-LIDARs spatial weighting" it was shown that the variance as measured with an upwind-looking (Spinner-based) Lidar, could be calculated from a low-pass Lorentzian-filtered turbulence of the Horizontal wave number spectrum  $F_u(k_1)$ 

$$\overline{\langle u \rangle^2} = \int_{-\infty}^{\infty} F_u(k_1) e^{-2z_R k_1} dk_1$$
(1.7)

That is, the lidar measured variance results from the Longitudinal turbulence spectrum low-pass filtered by an exponential filter with a cut-off wave number given by  $k_1 \approx 1/2z_R$ .

In standard constant azimuth ( $\varphi = 30^{\circ}$ ) LDA scanning mode, the ZephIR lidar measures a combination of the (*u*, *v*, *w*) velocity components. If we assume that the boundary layer turbulence is approximately isotropic on the limited scale of the Lorentzian filters HWHM parameter  $z_R$ , we can assume that the Lorentzian filter applies to all three velocity components, so we can define:

The Lorentzian optical probe volume is given by:

$$L_{Lorentzian}(k_1) = e^{-2z_R k_1}$$
(1.8)

# 2.2 Averaging associated with the three-revolution 3-s horizontal azimuth scans:

A simple model can be made if we assume that the resulting 3-s wind vector is obtained from an average of the stream wise wind component over the area covered by three revolutions:

An effective instantaneous horizontal averaging length scale can be estimated as the combined result of time lag and the circular coverage, which for the ZephIR lidar is equal to the scan diameter (The ZephIR lidar scans a horizontal circle of diameter *D* equal to  $\frac{1}{Cos(30)} = \frac{2}{\sqrt{3}}$  times the measurement height, i.e.  $(D \square 1.15z)$ . For example, at a measuring height of 100 meters, the lidar beam rotates at a speed of  $\pi D \square 363 [ms^{-1}]$ . With the ZephIR's inherent spectral sampling frequency of 200.000 samples per second, it corresponds to an azimuthal displacement of the laser beam of ~ 1.81 mm between two consecutive raw-spectral estimates. The ZephIR then averages such 4000 spectra during ( $5\mu$ S x 4000 averages), i.e. in ~ 20 milliseconds (50 Hz) over an azimuth distance corresponding to 1.81 mm times 4000scans equal to an azimuthal conical segment of ~ 7.26 meters.

In addition, the scan area covers a horizontal length scale given by the advection of the wind field by the mean wind speed during the ZephIR sampling time (3 s), see Fig.1:



Fig.1 The measurement area covered by the ZephIR lidar after three complete  $2\pi$ - azimuth scans, one per second, in a flow with mean wind speed U.

Therefore, a simple effective horizontal length scale,  $l_{az}$ , representing an effective filter-averaging length scale from sampling over three consecutive perpetual revolutions ( $6\pi$  azimuth), can to a first approximation be modelled by:

$$l_{az} = D + U\Delta T = \frac{1}{\cos(30)}z + 3U = \frac{2}{\sqrt{3}}z + 3U$$
(1.9)

If we model the effect of the 3-sec lasting 2 revolution azimuth scanning with a "Box car-like" filter function, we can further assume that the corresponding *Sinc* filter function  $(\frac{\sin^2 x}{x^2})$  applies as a low-pass filter on the turbulence in wave-number space, so that the combined 3-s  $6\pi$  azimuth filter function becomes

$$L_{Azimuth Scan}(k_{1}) = \frac{\sin^{2}(\pi k_{1}l_{az})}{(\pi k_{1}l_{az})^{2}}$$
(1.10)

### **3 The Effect on ZephIR Lidar measurements:**

With the above defined filters, we next investigate their combined effect on the QinetiQ ZephIR Lidar measured turbulence.

# 3.1 The combined Lorentz-filter and 3-s sampling effect on stream wise variance

For comparison with mast-mounted sonic anemometer measured variances, we first calculate the stream wise wind speed variance of the 3-sec averaged horizontal wind speeds, measured by the lidar (Lorentz- and Azimuth averaged), as :

$$\left\langle u^{2} \right\rangle_{ZephIR} = \int_{0}^{\infty} F_{u}\left(k_{1}\right) L_{A}\left(k_{1}\right) L_{L}\left(k_{1}\right) dk_{1}$$

with

$$L_{L}(k_{1}) = e^{-2z_{R}k_{1}}, L_{A}(k_{1}) = \frac{\sin^{2}(\pi k_{1}l_{az})}{(\pi k_{1}l_{az})^{2}}$$

where

$$F_{u}(k_{1}) = u_{*}^{2} \frac{z}{2\pi} \frac{102}{(1 + 33\frac{k_{1}z}{2\pi})^{5/3}} \left[\frac{m^{3}}{s^{2}}\right]$$

and

$$l_{as} = \frac{2}{\sqrt{3}} z + 3U \text{ and } z_R(z) \square 0.0012 \left(\frac{2}{\sqrt{3}} z\right)^2 = 0.0016 z^2 [m]$$

For the vertical wind speed profile we will assume typical danish Høvsøre Test Site parameters:

Roughness 
$$z_0 = 0.001[m]$$
; Friction velocity  $u_* = 0.5[ms^{-1}]; U(z) = \frac{u_*}{0.4} \ln \frac{z}{z_0} \Rightarrow U(100) \sim 15 ms^{-1}$ . (1.11)

The variance expression Eqs(1.11) is integrated in Mathematica, cf. Appendix I.



**Fig.2** Kaimal-modeled ZephIR stream wise wind speed standard deviations (inferred from consecutive three  $x 2\pi$  azimuth scans), relative to unfiltered (Sonic) values, as function of measurement height.



*Fig.3* Predicted ZephIR lateral wind speed standard deviations (inferred from consecutive three x 2  $\pi$  azimuth scans), relative to unfiltered (Sonic) values, as function of measurement height.



*Fig.4* Predicted ZephIR vertical wind speed standard deviations (inferred from consecutive three x 2  $\pi$  azimuth scans), relative to unfiltered (Sonic) values, as function of measurement height.
## 3.2 The filter effects on the ZephIR measured "TQE"

In the note "TQE & Shear stress\_Tensor\_from\_QQZephIR.doc", it was shown that the ZephIR lidar measured "turbulence parameter" could be compared to the standard expression for Turbulent Kinetic Energy, TKE, defined as

$$TKE = \frac{1}{2} \left( \langle u^2 \rangle + \langle v^2 \rangle + \langle w^2 \rangle \right)$$
(1.12)

but with the following modifications:

By use of 25% of the full  $\langle u^2 \rangle$  variance; 25% of the full  $\langle v^2 \rangle$  variance, and 150% of the full  $\langle w^2 \rangle$  variance, the QinetiQ ZephIR lidar's internal calculated "Turbulence parameter" was shown to be identical with a turbulence intensity based on the following definition of "Total "QinetiQ Eenergy":

$$TQE = \frac{1}{2} \left( \frac{1}{4} < u^2 > + \frac{1}{4} < v^2 > + \frac{3}{2} < w^2 > \right)$$
(1.13)

Based on the Kaimal spectra variance estimations in Eqs (1.2) we find

$$TKE \square 4.43u_*^2; TQE \square 2.06u_*^2, \text{ and } \frac{TQE}{TKE} \square 0.46$$
(1.14)

The definition of TQE in Eqs. (1.13) is defined in terms of "un-filtered" variances, that is, with no effects of the lidar's spatial and temporal averaging considered.

To investigate and evaluate the averaging effects of ZephIR measured turbulence, we next recalculate the variances in eqs. (1.13) including the filter effects.

Define the filter-averaged Total QinetiQ Energy in terms of ZephIR measured space and time averaged variances as

$$TQE_{av} \equiv \frac{1}{2} \left( \frac{1}{4} < u^2 >_{av} + \frac{1}{4} < v^2 >_{av} + \frac{3}{2} < w^2 >_{av} \right)$$
(1.15)

Where we as above calculate the ZephIR averaged variances by filtering, viz.:

$$\langle u \rangle_{av}^{2} = \int_{0}^{\infty} F_{u}(k_{1}) L_{A}(k_{1}) L_{L}(k_{1}) dk_{1}$$

$$\langle v \rangle_{av}^{2} = \int_{0}^{\infty} F_{v}(k_{1}) L_{A}(k_{1}) L_{L}(k_{1}) dk_{1}$$

$$\langle w \rangle_{av}^{2} = \int_{0}^{\infty} F_{w}(k_{1}) L_{A}(k_{1}) L_{L}(k_{1}) dk_{1}$$

$$\text{where: } L_{L}(k_{1}) = e^{-2z_{R}k_{1}}; L_{A}(k_{1}) = \frac{\sin^{2}(\pi k_{1}l_{az})}{(\pi k_{1}l_{az})^{2}}$$

and with standard Káimál power spectra (in wavenumber presentation) given by

$$F_{u}(k_{1}) = u_{*}^{2} \frac{z}{2\pi} \frac{102}{(1 + 33\frac{k_{1}z}{2\pi})^{5/3}} \left[\frac{m^{3}}{s^{2}}\right]$$

$$F_{v}(k_{1}) = u_{*}^{2} \frac{z}{2\pi} \frac{17}{(1 + 9.5\frac{k_{1}z}{2\pi})^{5/3}} \left[\frac{m^{3}}{s^{2}}\right]$$

$$F_{w}(k_{1}) = u_{*}^{2} \frac{z}{2\pi} \frac{2.1}{1 + 5.3\left(\frac{k_{1}z}{2\pi}\right)^{5/3}} \left[\frac{m^{3}}{s^{2}}\right]$$
(1.16)

As before: 
$$l_{as} = \frac{2}{\sqrt{3}} z + 3U(z)$$
 and  $z_R(z) \square 0.0016 z^2 [m]$ .



*Fig3. Prediction of ZephIR sampled TQE turbulence relative to unfiltered (Sonic) variance TKE, as function of measurement height. (Averaging time corresponding to 10-min).* 



*Fig4. Prediction of 3-sec averaged TQE turbulence relative to unfiltered (Sonic) variance TKE, as function of measurement height.* 



Fig5. Prediction of 3-sec averaged  $TQE_{3s}$  turbulence, relative to 3-s averaged (Sonic) turbulence  $TKE_{3s}$  as function of measurement height.

## Appendix I:

Mathematica filter evaluation program: lidarfilter\_HEJ\_TM\_04.nb

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Deve	lopment of LIDAR wind sensing for the German offshore tes	t field 6
	Objectives of the LIDAR r	esearch proposal
·	<ol> <li>Development and demonstration in</li> <li>Power curve measurements withou Offshore capability of the LIDAR sy (in preparation for measurements ii inflow of Multi-MW wind turbines</li> <li>Development of wind field simulation farm operation including dynamic w</li> <li>Measurements of turbulence properesolution as base for new and fast determination</li> </ol>	four typical areas: t met mast stem h the offshore test field) dds in dynamic wakes and in the on techniques for inflow in wind wake effects rties of windfields in a high er methods for power curve
•	Formulation of standardised power offshore test field taking into consi guideline "Part 2: Determining the Standardised Energy Yields "1)	<sup>•</sup> curve measurements in the deration the FGW technical Power Performance and
•	Provision of LIDAR hardware and c application in the offshore test field	of the know-how needed for the I and other R&D projects
1): [	ttp://www.wind-fgw.de/tr engl.htm	
sw	E Stiftungslehrstuhl Windenergie am Institut für Flugzeugbau	Universität Stuttgart 🎆





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70569 Stuttgart, Germany	
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Variable	Symbol	Lidar Perspective	
Mixed layer	k	Strength of back-scatter	
height (Boundary		signal identifies aerosol	<ul> <li>List of Boundary</li> </ul>
layer depth)		layer(s). (Mok and	Laver Parameters
		Rudowicz, 2004)	Layer r diameters
Mean velocity	₩ [u,v]		<ul> <li>obtained from</li> </ul>
Turbulence	σ.,	From w'	lidar data
Eddy dissipation rate	٤	From spectra <i>s</i>	nuar uata
Lagrangian	•	Decay tim e scales for auto	<ul> <li>for dispersion</li> </ul>
integral timescale	$\tau_{\star} = \int R(\tau) d\tau$	correlation coefficient	modelling
and	1.1 1	$\frac{1}{1}$ $\frac{1}$	mouening
Integral length		$R(\tau) = \frac{\alpha (\tau) \alpha (\tau + \tau)}{\tau} \text{ for }$	
scale	-	$\sigma_{\bullet}$	
	$L_{i} = \int R(s) ds$	lag Γ	
	ō	$R(s) = \frac{u'(x)u'(x+s)}{\sigma_{\pi}^2} \text{ for }$	<ul> <li>Helsinki data</li> </ul>
		lags.	primarily vertically
Sensible heat flux	Horg	Indirectly from lidar third	pointing
	$H = Q_H$	m om ent $\overline{\psi'}^3$ . (Gal-Chen et	
Flux of	$= \rho C_{-} \overline{\psi' \theta'}$	al 1992)	
temperature fluctuation	₩'&'	Orfrom weas below.	<ul> <li>Collier et al 2005 (Bulletin of AMS)</li> </ul>
Convective		From $\sigma^2 \approx \beta w^2$	
velocity scaling	[ [μ' θ']] <sup>3</sup>	(angraging at al 1004)	
(for unstable conditions).	$w_{\bullet} = \begin{bmatrix} ng \\ \theta \end{bmatrix}$	where β ? 0.52 within 0.2< z/k<0.5	
	1		

## Autonomous Doppler lidar system - requirements

Traditionally lidar systems have been expensive, bulky and require a high level of routine maintenance.

1998 – 2006:  $CO_2$  TEA system, 112 m range gates, 600 m min. range, housed in 4.6 tonne vehicle, dedicated operator



Doppler lidar system designed and developed to meet the following objectives:

Long term velocity and backscatter measurements in urban and rural environments – system is eye safe, near IR

✤ Require measurements from close to street level → top of boundary layer – high spatial resolution to retrieve measurement at many levels in urban canyon

\* Air quality and pollution dispersion – high temporal resolution of system to retrieve turbulence and velocity variance information

Field deployable – compact system

\* No dedicated user required – remote access to system with software that will allow full control of system, view and download data

## *dar* Autonomous Doppler lidar system for range resolved remote sensing of the atmosphere

Parameter	Value
Operating wavelength	1.5 microns
Pulse repetition frequency	20 kHz
Beam divergence	50 μrad ~ 5 cm at 1 km
Range gate	20 – 60 m
Minimum range	30 m
Maximum range	Up to 7 km
Temporal resolution	0.1 – 30 s
Optical base unit (1) Antenna (2) Signal processing and data acquisition unit (3)	56 x 54 x 18 cm 8 cm diameter Standard desktop pc
Hemispheric scanner & Video camera for alignment and sighting	50 mm aperture 0 – 360 ° azimuth 0 – 180 ° elevation 0.5 ° resolution




























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### Doppler Lidar Measurements Using a Fibre Optic System

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## Performance in Practice

- Reliable measurements at optimal range for visibility <40 km.</li>
- Returns are marginal from ranges where the bistatic optics are not optimal, particularly when the optics are wet.
- Polarization scrambling requires a 1 min sampling time at each range, implying a long cycle time around all ranges.



Signal voltage for path delay of 
$$\tau$$
:  

$$V(t,\tau) = V_0 \cos(\phi_0 + \Delta \phi(t,\tau) + 2\pi f_B t),$$
Autocorrelation function :  

$$R(t,t',\tau) = V_0^2 \cos(\phi + \Delta \phi(t,\tau) + 2\pi f_B t) \cdot \cos(\phi_0 + \Delta \phi(t',\tau) + 2\pi f_B t')$$

$$= \frac{1}{2} V_0^2 \{\cos(\Delta \phi(t,\tau) - \Delta \phi(t',\tau) + 2\pi f_B \overline{t-t'}) + \cos(2\phi_0 + \Delta \phi(t,\tau) + \Delta \phi(t',\tau) + 2\pi f_B \overline{t+t'})\}.$$

Orthogonality ...  $< \cos(\Delta\phi(t,\tau') - \Delta\phi(t',\tau')) \cdot R(t,t',\tau) >$   $= (V_0^2/4) \cos(2\pi f_B \overline{t-t'}), |\tau - \tau'| < \tau_c,$   $\approx 0, otherwise.$ where  $\Delta\phi(t,\tau) = \phi(t) - \phi(t-\tau)$ So we don't have to measure the absolute phase: we only have to monitor its drift with time



Mathematically,  $V_{a}(t, \varepsilon) = V_{a, \max} \cos(\Delta \phi(t, \varepsilon))$   $V_{r}(t, \varepsilon) = V_{r, \max} \sin(\Delta \phi(t, \varepsilon))$ . where  $\varepsilon$  is the delay around the loop. Hence  $\Delta \phi(t, \varepsilon) = \tan^{-1}(V_{r}/V_{r, \max}, V_{a}/V_{a, \max})$ and  $\Delta \phi(t, \tau) = \sum_{k=0}^{k < \tau/\varepsilon} \Delta \phi(t - k\varepsilon, \varepsilon)$ .















































- WindCube is a pulsed Lidar :
  - Simultaneous measurement at any height
  - Steady performances whatever the height
- WindCube is upgradable:
  - 600m detection
  - 3D windflows
- WindCube is adapatable to reach higher ranges
- WindCube has a 15° scaning angle
- WindCube is robust













Summary of IEA RD&D Wind – 51<sup>st</sup> Topical Expert Meeting on

### State of the art of Remote Wind Speed Sensing Techniques using Sodar, Lidar and Satellites

January 2007, Risø, Denmark

### Background

Wind power is moving towards the installation of wind farms in complex terrains, offshore, in forests, and at high levels in the atmosphere. Marketing of large, multi-MW wind turbines is in continued growth. At the same time our basic knowledge on winds in these challenging environments is inadequate.

The method traditionally used for accredited measurements for wind energy purposes is to mount cup anemometers on met masts. As turbines grow in height, mast instrumentation, erection and maintenance have become expensive; prices increase with height and building permits can be time-consuming. At the same time the discrepancies between the measured wind at the rotor centre and the turbine performance have increased the need for determining the wind over the whole turbine rotor.

Successful development of wind power should be based on sound information on winds in each location. To achieve this it is important to place emphasis on new observation methods and strategies. Most promising are the new (for wind energy purposes) remote sensing techniques: Sodar, Lidar and satellite. Sodar is based on sound propagation, Lidar on laser doppler and satellite on microwave scatterometry and Synthetic Aperture Radar (SAR) methods. Advantages and limitations of the various techniques will be described and discussed.

### SODAR

Sodar (SOund Detection And Ranging) provides a method for wind speed measurements. The instrument is ground-based and emits a short pulse of sound at a certain frequency to the atmosphere. The sound propagates upwards, while at the same time a part of the sound is reflected back. The Doppler frequency shift of the received signal is proportional to the wind speed aligned to the transmission sound path. By combining three or five of these pulses, usually one along the vertical and two or four inclined to the vertical, the three-dimensional velocity field of both the mean values and the turbulent values is calculated.

### LIDAR

Lidar is a remote sensing technique that offers the ability to determine wind speed and direction at substantial heights using a ground-based instrument. In this respect it is similar to Sodar, but operates via the transmission and detection of light rather than sound. The basic Lidar principle is to measure the Doppler shift of radiation scattered by

natural aerosols carried by the wind. Typically, these are dust, water droplets, pollution, pollen or salt crystals. A new generation of fibre-based Lidar has emerged in recent years that operates close to the theoretical limit of sensitivity and typically only needs to detect one photon for every 10E+12 transmitted in order to measure wind speed. Since the Doppler-shifted frequency is directly proportional to line-of-sight velocity, the wind speeds obtained by a Lidar instrument seem not to need calibration. This however still remains to be documented by more measurements and by a full description of the whole measurement chain. As in the case of Sodar, the Lidar is also a new instrument, and its merits and limitations are neither fully documented nor known. In the case of the Lidar, the measurement of the wind speed takes place on the surface of a cone where the depth changes as a function of the focus distance. The measurement of the turbulence quantities using Lidar also remains to be documented.

### Satellite remote sensing

Satellite remote sensing provides wind maps (snap-shot images) of the surface wind at 10 m above sea level. From a scatterometer, twice daily, wind maps at grid resolution of 25 km are available. The data series from July 1999 to present holds more than 5000 observations at most locations of the globe. Due to the resolution of 25 km, observations are not available close to the coastline (usually there is a void around 40 to 50 km distance offshore). In contrast, SAR wind maps cover the near coastal zone in which most wind farms are located. Far fewer SAR wind maps are available (e.g. a few hundred or less), but by using statistical treatment of a few samples, rough estimates of the wind resource can be obtained. The accuracy, around 1.1 m/s standard error, on a series of wind maps compared to offshore mast observations is useful in pre-feasibility studies and in decisions about the location of offshore masts (or LIDAR/SODAR). In addition, if high-quality met-observations are available within a mapped area, the relative differences in winds between different locations can be estimated with higher accuracy, possibly around 0.6 m/s.

### **Participants / Presentations**

A total of 51 participants attended this meeting with representatives from Denmark, Finland, Germany, Ireland, Norway, Sweden, the Netherlands, UK and USA. The participants mainly represented National Research Organizations, utilities and entities performing measurements.

The large number of participants in the meeting reflected the interest in this research topic and application in wind turbine work. The number of participants was restricted due to the size limitations of the meeting facilities.

The number of presentations was 29, covering the following subjects:

General	8 presentations
Sodar	10 presentations
Lidar	9 presentations
Satellite	2 presentations

### Discussion

A discussion was held at the end of the meeting. Some of the discussions are summarized below. These points should not be regarded as "truths" coming out of the discussions, but rather comments that participants gave.

General

- There was a common understanding that there is a need for more experience from remote sensing, especially comparing the performances of Lidar and Sodar.
- Lidar and Sodar will complement each other for a while. Both instruments will have a future in atmospheric science.
- Axel Albers: Both Sodar and Lidar have room for improvements. I researched Sodar since 1992. We never got the reproducibility we now see with the Lidar. The first QinetiQ Lidar give astonishing results. It will take a very long time before Sodar can replace met mast in terms of absolute wind speed. This will soon happen with Lidar.
- Andrew Tindal: For some time to come remote sensing will be used in conjunction with conventional anemometry. But, carefully, we should step towards the replacement, through understanding all the errors.

Sodar

- Sodar are commercially available from a number of different companies. Lidar on the other hand are for sale, but are not as developed and commercialized.
- Sodar are generally cheaper than Lidar. A price tag of the ZephIR is 100.000 GBP. Axel Albers commented that customers asking for measurements are not willing to pay rental for such expensive instruments.
- Sodar has fundamental limitations compared to Lidars. The wave length of the sound is large compared to that of light, implying bulkier sodar instruments. The speed of sound is much smaller than that of light, implying that the sound ray propagation in the atmosphere is considerably more complicated, e.g. beam drift. Given the recent development some argued that Lidar has a brighter future than that of Sodar.

Lidar

- Lidar has the disadvantage that the averaging volume increases with height, whereas the corresponding volume for the Sodar remains constant with height. Maybe the pulsed lidar technology will change that.
- Hans E. Jørgensen pointed out that we need to test the performance of Lidar in complex terrain: wind shear, turbulence intensity and flow inclination are issues here of great interest for developers.
- Troels Friis Pedersen: I believe a Lidar mounted on nacelles will be extremely useful for power performance measurements. Stefan Emeis: Maybe there is a difference in the needed accuracy between siting and power performance measurements. Sodar may be fine for wind profiles. J. Højstrup strongly

disagreed. We always need the same accuracy. Better accuracy implies lower financial and technical uncertainties. Albers: There are still a lot of uncertainties in site assessment.

Satellites

- Satellites always see the structure of the surface, e.g. SAR see the wind stress on the surface. Models are needed to transfer this information to hub height. Given the accuracy needed it may not be worthwhile.
- Space-borne Lidar are coming and they may be useful.
- Neil Douglas (Natural Power Consultants): Maybe accuracy is not always so important. For example satellites may be used for relative resource estimation.

There is a need for "best practices" on how to use remote sensing as siting devices, etc., as suggested by Kathleen Moore. More sodar /lidar/mast comparison needs to go to the literature. The initiative of Risø of a remote sensing test facility at Høvsøre is good!

### Continuation

There was a common understanding that there is a need for more experience from remote sensing in order to increase the accuracy and the repeatability of measurements, especially comparing the performance of Lidar and Sodar with anemometers. The IEA-developed Recommended Practices for anemometry are available and could be used as a reference for developing similar documents for Lidar and Sodar. Participants pointed out that such documents are needed in a near-future time frame.

As a first step of continuation it was considered relevant to undertake initial work related to develop such practices. It was agreed to form two Ad-Hoc groups to put together proposals for the proper operation of a Sodar/Lidar. The ad-hoc groups should make to-do lists for improvements of the instruments.

- Sodar group: Kathleen Moore will take the lead. Participants: Gunter Warmbier, Mats Hurtig, Andy Oldroyd, Finn Nyhammer, Brian Hurley, Peter Clive, Sabine vonHunerbein, Ken Underwood, Stuart Bradley
- Lidar group: Ioannis Antoniou will take the lead, Axel Albers, Ian Locker, Detlef Kindler, Andreas Rettelmeyer, Brian Hurley

It was noted that there exists a general recommended practice for remote sensing. One in Germany (VDI 3786 Part 14, Verein Deutscher Ingenieure, Environmental meteorology, Ground-based remote sensing of the wind vector. Doppler Wind LIDAR, Dec. 2001) and elsewhere.

The results from the Ad-Hoc groups will be reported at the upcoming meeting of the IEA Wind Executive Committee by the Operating Agent of Task 11. This may result in further action within this field.

## List of participants

# IEA R&D Wind Task11, Topical Expert Meeting 51 State of the art of Remote Wind Speed Sensing Techniques using Sodar, lidar and Satellites Risô, Roskilde, Denmark 23-24 January 2007

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